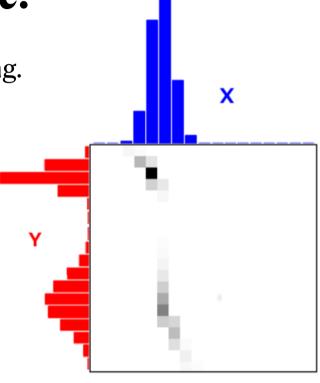
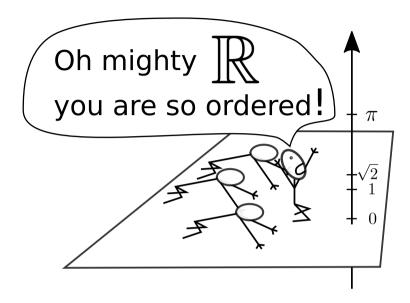
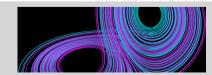
An introduction, with application to climate modelling.

(joint with Mat Chantry, Milan Klöwer & Tim Palmer)

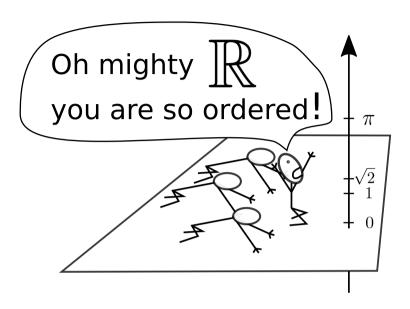


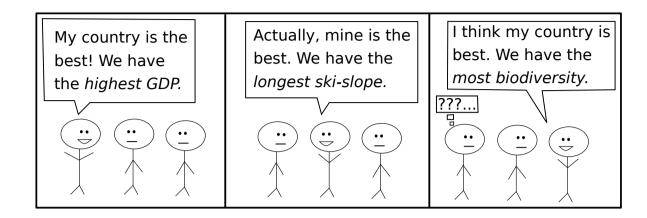




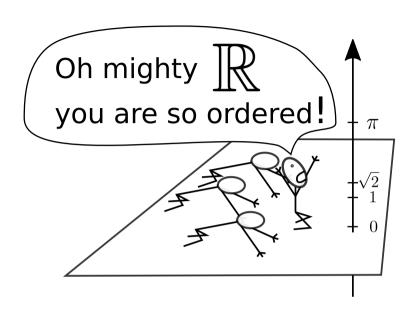


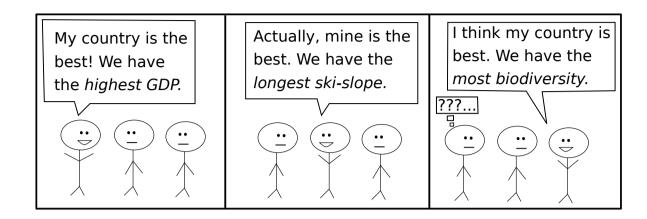






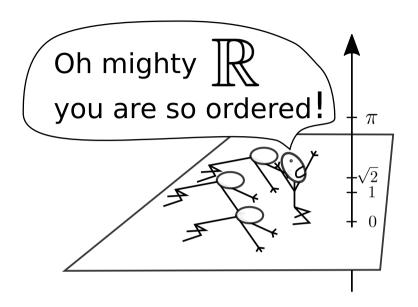


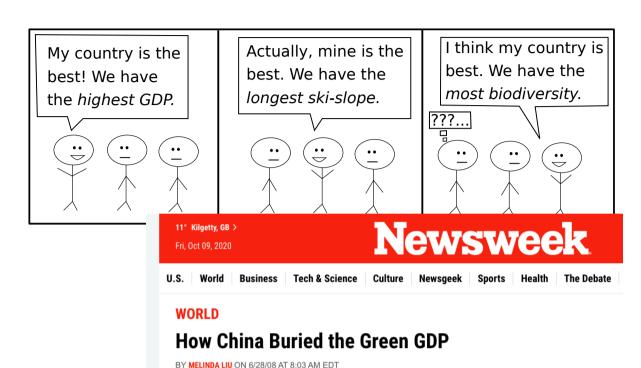




·Real world problems are multi-dimensional.





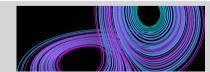


SHARE f y t in p 6

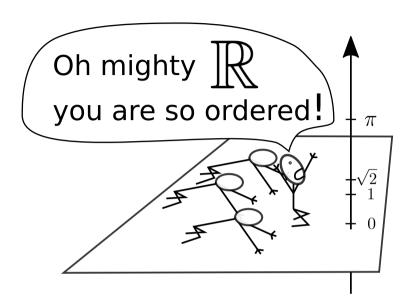
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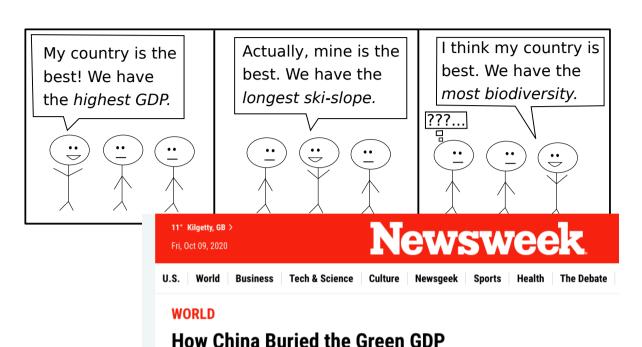
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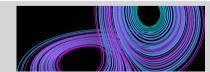
BY MELINDA LIU ON 6/28/08 AT 8:03 AM EDT

SHARE f y t in p 6 2

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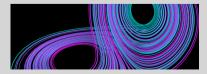
•If anyone says that their metric is the *best*, you should probably be cynical!

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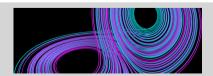
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Plan of talk:

- 1. What is the Wasserstein distance?
- 2. What are the advantages of the WD, and how to compute it.
- 3. An application: exploring model climatology in low-precision.









•The WD (Earth Mover's distance) is a distance between probability distributions (measures) $\mu \& \nu$.



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Predictability group internal seminar 09.11.20

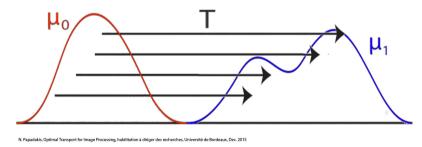




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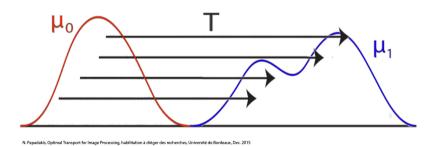




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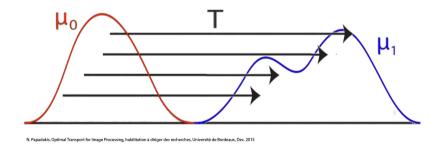


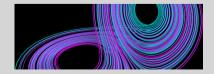


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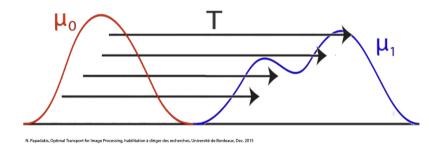




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- •For the case $c(x, y) = |x y|^p$ we call the optimal cost the p-Wasserstein Distance (we'll always take p = 1)





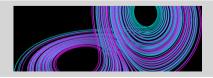








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• A transport strategy is a permutation of N objects $\sigma \in S_N$.

The cost of a strategy is
$$\frac{1}{N} \sum_{i=1}^{N} c(x_i, y_{\sigma(i)})$$
.



$$WD_1(\mu, \nu) := \min_{\sigma \in S_N} \frac{1}{N} \sum_{i=1}^N |x_i - y_{\sigma(i)}|$$









Suppose

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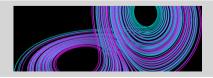
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nb. when $M_1 = M_2 = N$ and $p_i = q_i = \frac{1}{N}$ it turns out the two definitions are equivalent.









- 2) What are the advantages of the WD?
 - (i) It metrizes the space of probability distributions.





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If μ_k is a sequence of probability distributions, then

$$WD_1(\mu_k, \mu) \to 0$$
 if \mathcal{E} only if $\mu_k \to \mu$ (weak \star)

where $\mu_k \to \mu$ (weak \star) means:

$$\int_{\mathbb{R}^n} \phi(x) d\mu_k(x) \to \int_{\mathbb{R}^n} \phi(x) d\mu(x) \text{ for any bounded function } \phi(x)$$



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nb. (i) \Longrightarrow It takes into account the whole distribution (i.e. "all moments")



(ii) It is versatile.





(ii) It is versatile.

You can compare *any* two probability distributions:

- · Continuous distributions.
- Discrete / singular distributions.
- Distributions defined on different spaces.

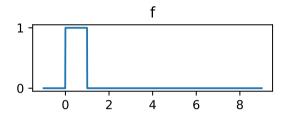


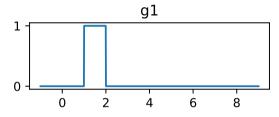


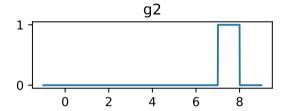




•Consider the following 3 simple PDFs:

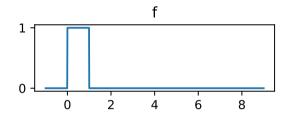


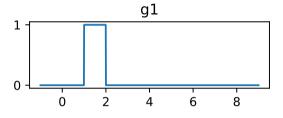


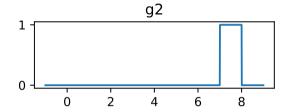




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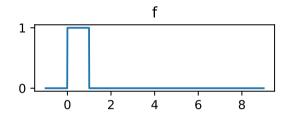


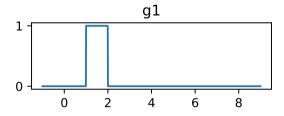
With L^p -distance we have $||f - g_1||_{L^p} = ||f - g_2||_{L^p} = 2$.

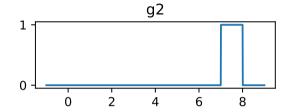
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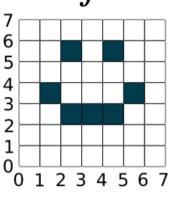


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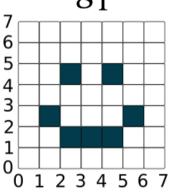
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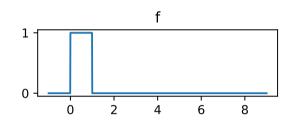


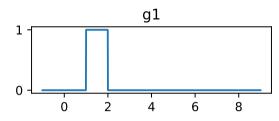


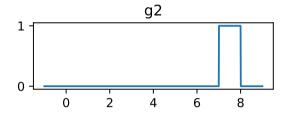


Nb. This is a shortcoming of many common metrics e.g. K-S test / K-L divergence

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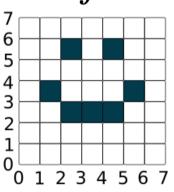


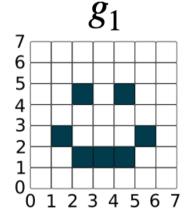
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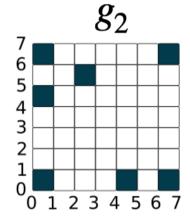
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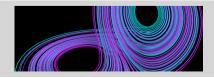




E. Adam Paxton









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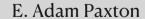


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All of these can be found at github.com/eapax/EarthMover.jl











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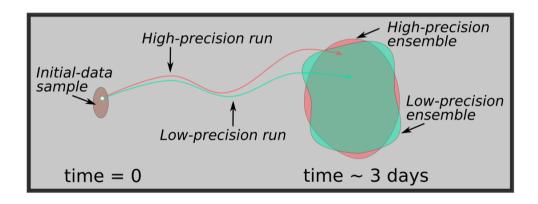
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- •As forecast models move to low-precision, it's natural to ask if these models are suitable for climate modelling (some have argued NOT).





Climate modelling & weather forecasting are different methodologies.

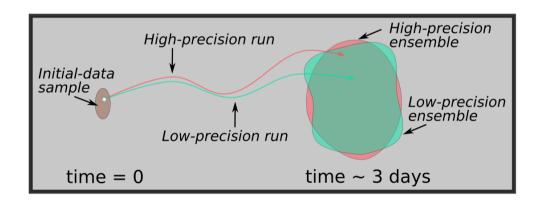
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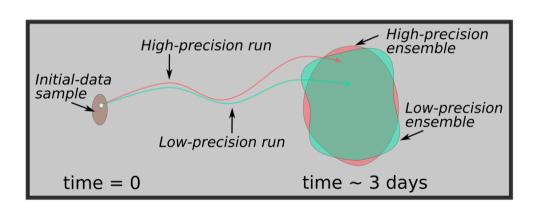






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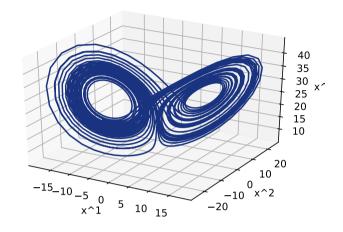
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Idea: use the Wasserstein Distance to test this.





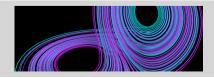


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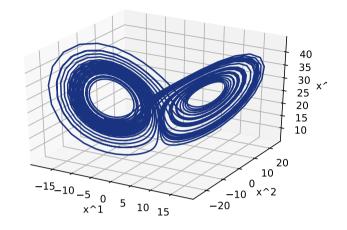
$$\dot{x}^{2} = \left(\frac{8}{3} - x^{3}\right) x^{1} - x^{2}$$

$$\dot{x}^{3} = x^{1}x^{2} - 28x^{3}$$





•Admits an attractor $\mathscr{A} \subseteq \mathbb{R}^3 \ (x(t) \to \mathscr{A} \text{ as } t \to \infty).$



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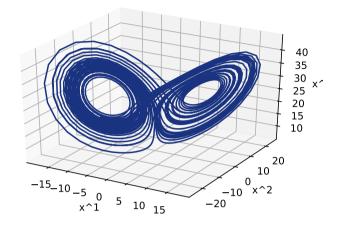
$$\dot{x}^{2} = \left(\frac{8}{3} - x^{3}\right) x^{1} - x^{2}$$

$$\dot{x}^{3} = x^{1}x^{2} - 28x^{3}$$





- •Admits an attractor $\mathscr{A} \subseteq \mathbb{R}^3 \ (x(t) \to \mathscr{A} \text{ as } t \to \infty).$
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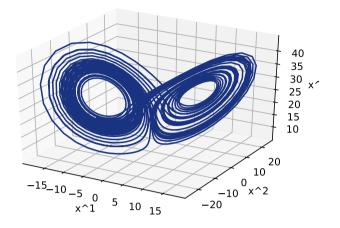




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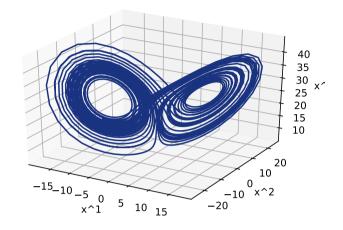
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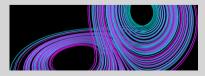


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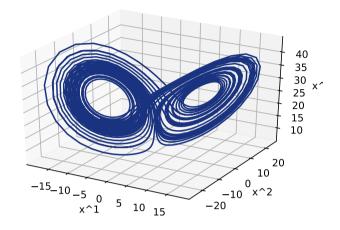


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nb. link to weak★ *convergence!*



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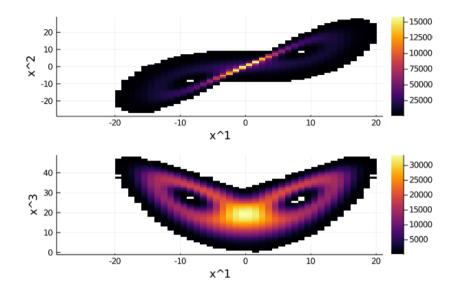
Two methods:





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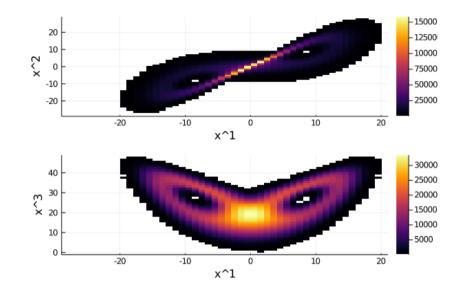
1. Data-binning (i.e. approximate μ as a histogram)

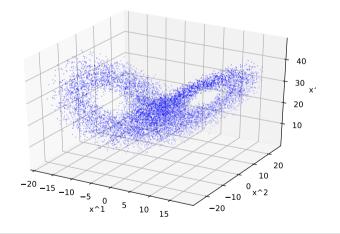




Two methods:

- 1. Data-binning (i.e. approximate μ as a histogram)
- 2. Scatter-plotting (i.e. approximate directly from sampling as $\mu \approx \frac{1}{N} \sum_{i=1}^{N} \delta_{x_i}$)





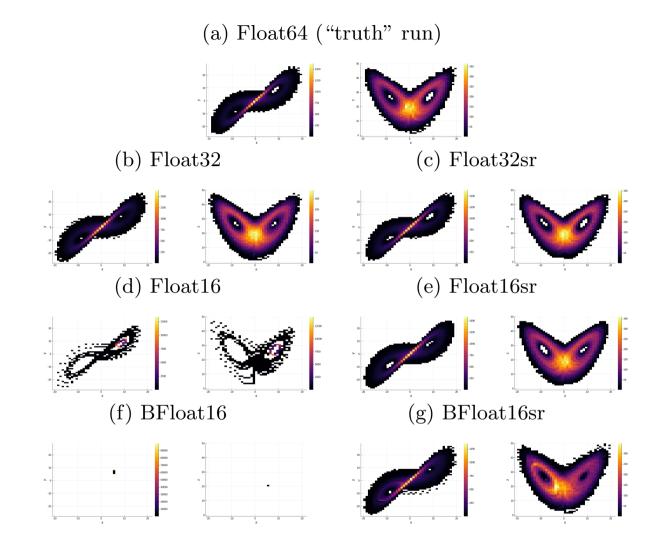






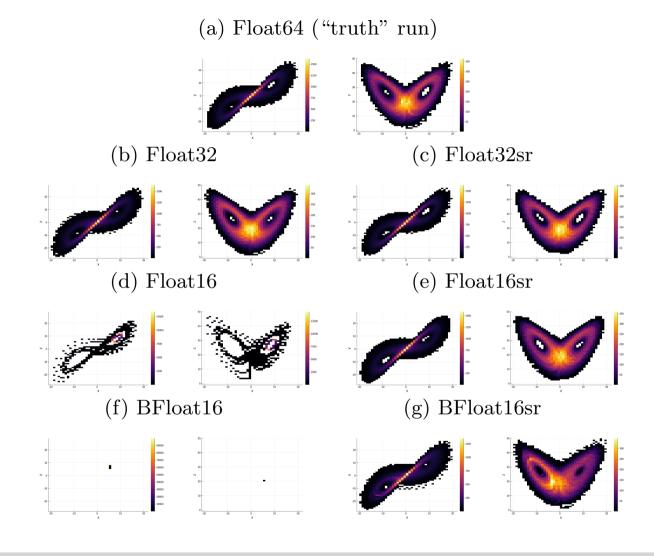


•Integrated L63 in different numerical precisions.



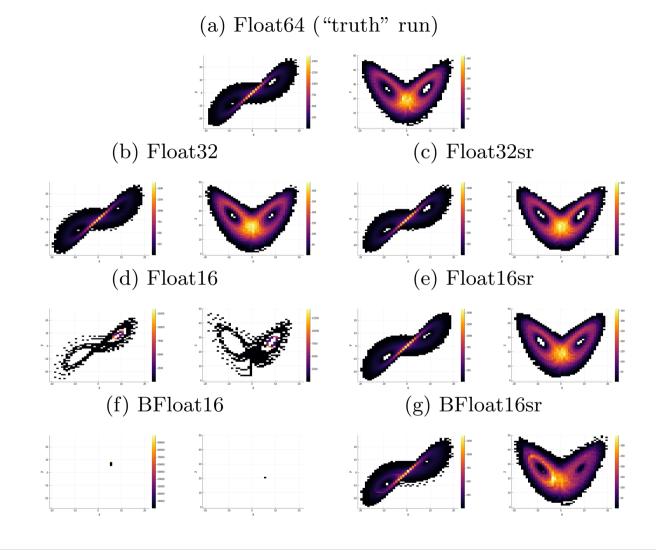


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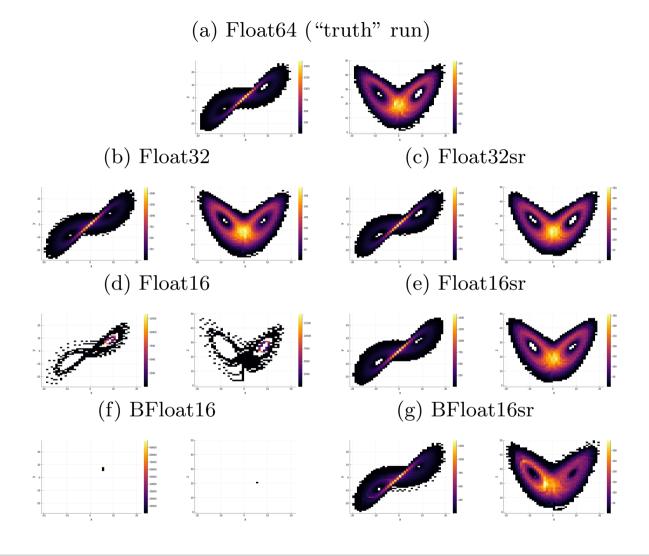


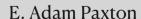






- •Integrated L63 in different numerical precisions.
- Approximated invariant measures by data-binning.
- •We want a method for quantitative comparison.
- •Let's compute the Wasserstein Distances!









precision	WD(precision, Float64)
Float64	0.0
Float32	0.456
Float32sr	0.353
Float16	14.8
Float16sr	0.421
BFloat16	16.1
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·We need a null hypothesis.

·Idea: use an ensemble.

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Experiment set-up:

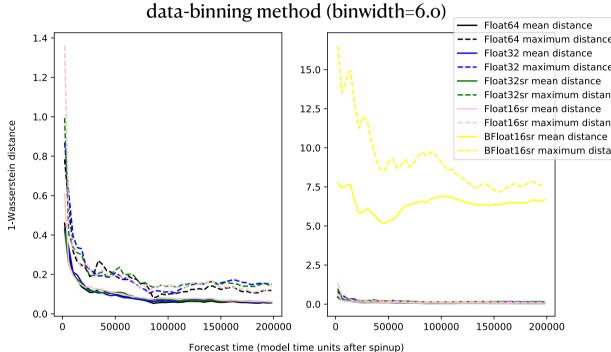
- Take one 5-member Float64 ensemble (Control)
- •Take a 5-member ensemble for each precision (including Float64) and compare with the Control pairwise (25 comparisons).
- •Plot the mean & maximum values with time.



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Convergence to statistical equilibrium:







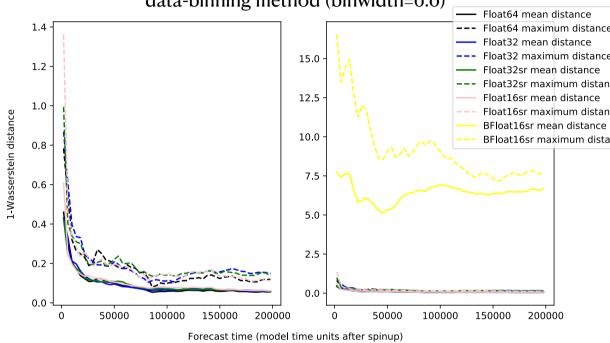


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The Float64 vs Control test (black lines) serves 2 purposes:

Convergence to statistical equilibrium: data-binning method (binwidth=6.0)



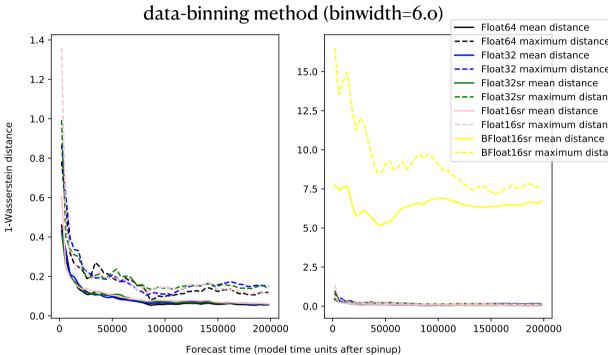
- 1. It gives a null hypothesis.
- 2. It shows that enough time has elapsed to reach statistical equilibrium.

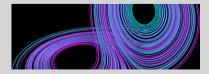






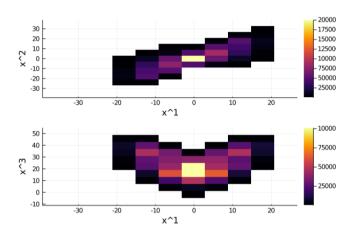
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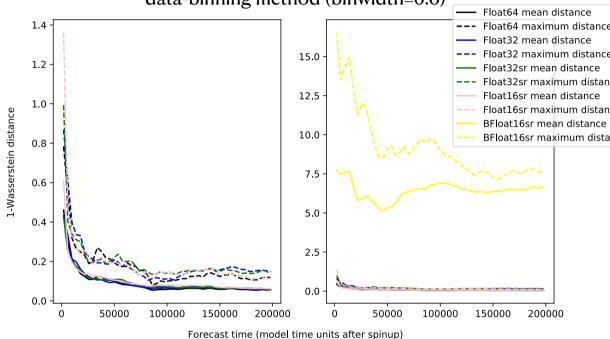


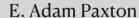


nb. bin-width=6.0 looks like:



Convergence to statistical equilibrium: data-binning method (binwidth=6.0)



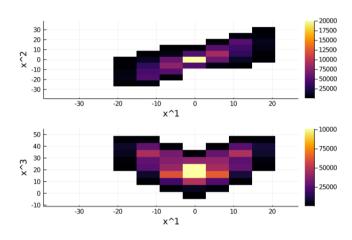






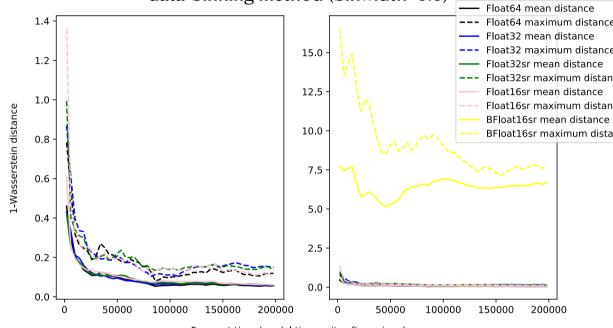


nb. bin-width=6.0 looks like:



•Results are not sensitive to decreasing bin-width.

Convergence to statistical equilibrium: data-binning method (binwidth=6.0)



Forecast time (model time units after spinup)





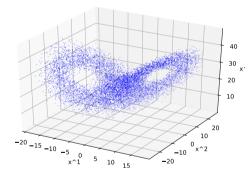
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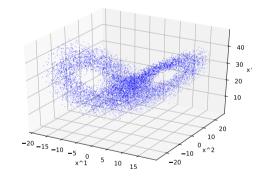




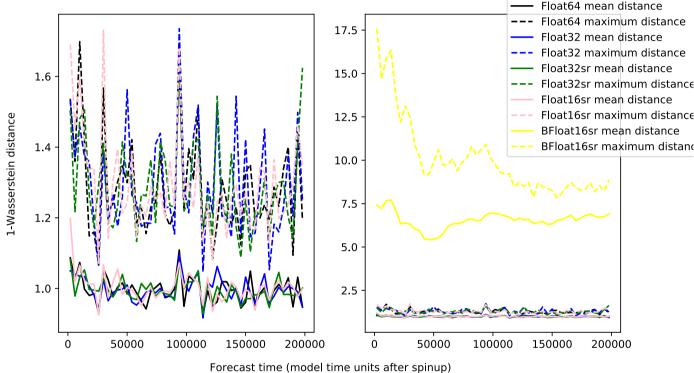
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(i.e. approximate as

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).



Convergence to statistical equilibrium: scatter-plot method (sample size=2500)



It gives comparable results.

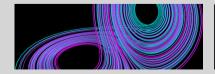
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Shallow Water Model:

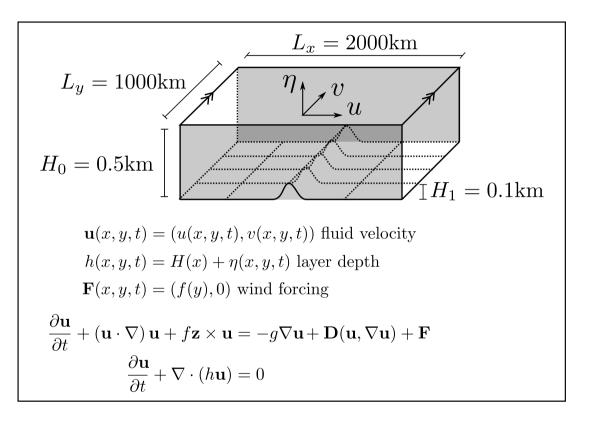
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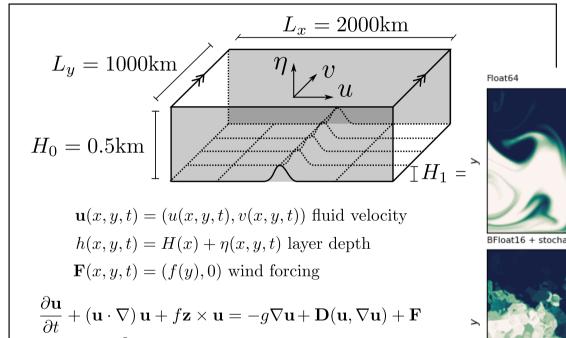


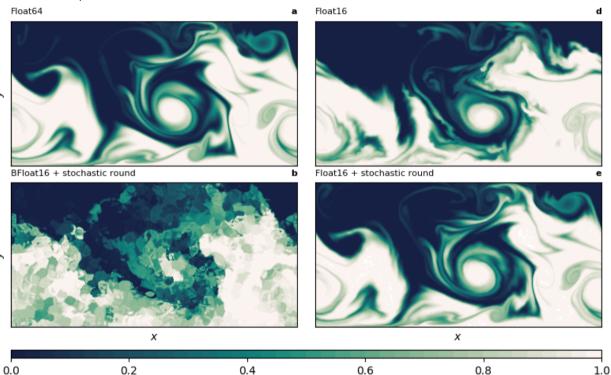


Shallow Water Model:

github.com/milankl/ShallowWaters.jl

•Finite difference scheme, 100×50 spatial grid



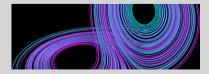


Tracer concentration



 $\frac{\partial \mathbf{u}}{\partial t} + \nabla \cdot (h\mathbf{u}) = 0$

We want to estimate the Shallow Water model climatology (i.e. invariant measure).





We want to estimate the Shallow Water model climatology (i.e. invariant measure). Some problems arise:

- ·We have time evolution in a $100 \times 50 = 5000$ dimensional space.
- · Working with high-dimensional probability distributions is non-trivial.
- •Data-binning becomes stupid. Looking at just one parameter u and assigning just 2 bins per spatial coordinate would lead to 2^{5000} bins. (number of atoms in observable universe $\approx 2^{270}$)



•One strategy: project down onto lower-dimensional subspaces.



- One strategy: project down onto lowerdimensional subspaces.
- •This is what I have seen done so far.

Ranking IPCC Models Using the Wasserstein Distance

G. Vissio¹, V. Lembo¹, V. Lucarini^{1,2,3}and M. Ghil^{4,5}

¹CEN, Meteorological Institute, University of Hamburg, Hamburg, Germany
²Department of Mathematics and Statistics, University of Reading, Reading, UK
³Centre for the Mathematics of Planet Earth, University of Reading, Reading, UK
⁴Geosciences Department and Laboratoire de Météorologie Dynamique (CNRS and IPSL),
Ecole Normale Supérieure and PSL University, Paris, France
⁵Department of Atmospheric & Oceanic Sciences, University of California at Los Angeles,
Los Angeles, USA

Key Points:

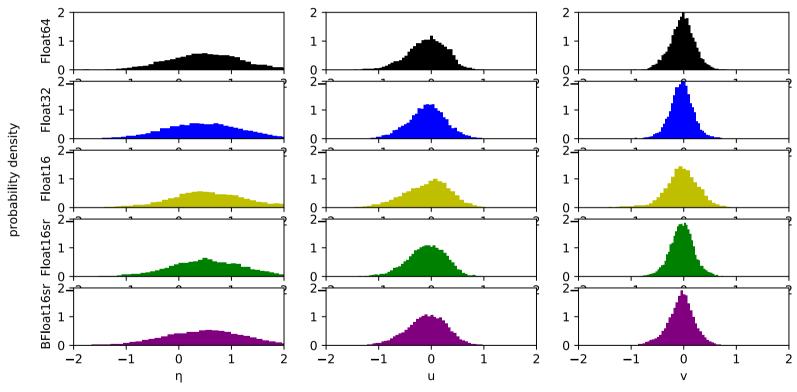
- \bullet Evaluation of climate model performance by benchmarking with reference datasets
- \bullet Climate model ranking related to the choice of variables of interest
- · Highlighting model deficiencies through emphasis on climatic regions and variables





We can do this for Shallow Waters. Take spatial average over some (arbitrary) region (100,200)km x (400,500). Do 1D data-binning.

Model climatology for Shallow Water Equations after single 20 year run (values spatially averaged over (lon,lat)∈(100, 200)km x (400, 500)km region)

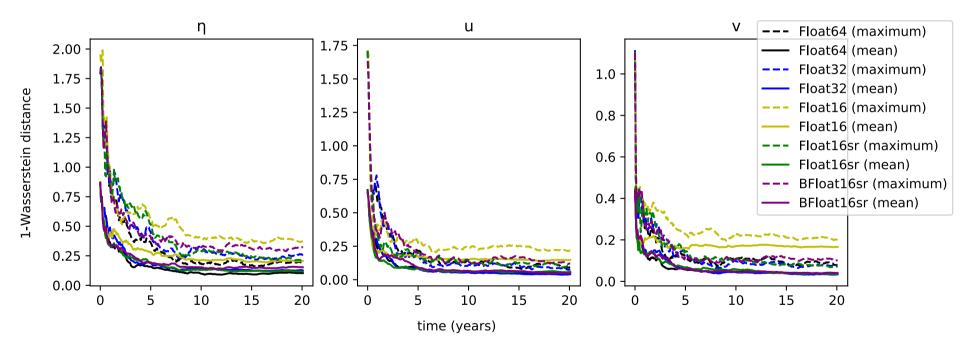






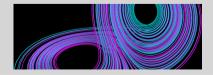
- · We can compute Wasserstein distances between these 1D distributions.
- Same experiment as before (5-member ensembles, one Control ensemble).

Shallow Water Equations: Convergence of Wasserstein distances (data-binning method, values spatially averaged over (lon,lat)∈(100, 200)km x (400, 500)km region)



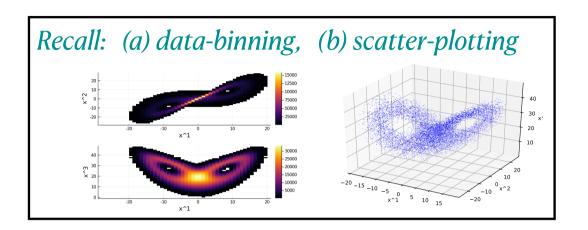


• The problem with projection is you are no longer considering the full distribution.





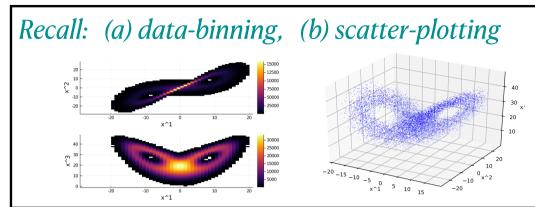
- •The problem with projection is you are no longer considering the full distribution.
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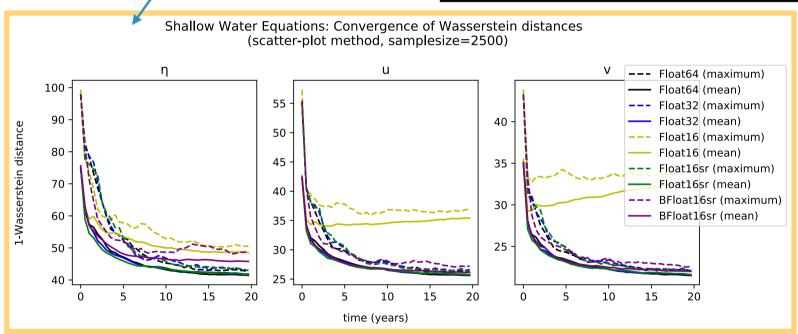


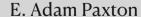


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 This seems to work!!!











Conclusion of experiment.

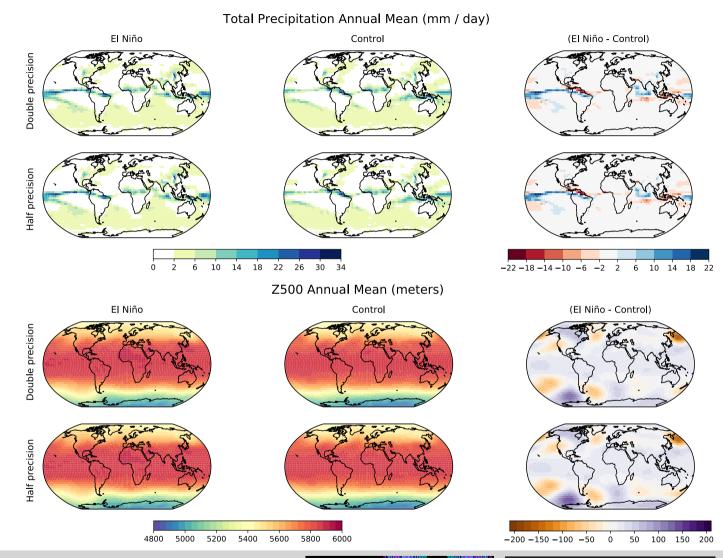
The results provide strong evidence that the effects of rounding error on the shallow water model climatology, when compared with initial condition variability & discretisation error are:

- 1. Negligible for **Float32** and **Float16sr**.
- 2. Significant for **Float16** and **BFloat16sr**.





- Next steps: performing the same analysis to reduced precision SPEEDY.
- A coarse resolution (3.75° × 3.75°) atmosphere only, primitive equation model (prescribed SSTs) with simplified parameterisations.
- Leo's 16-bit (deterministic) version of the code has held up to the first tests.



E. Adam Paxton





Summary of talk:

- The Wasserstein metric gives a notion of distance between probability distributions.
- It has excellent properties.
- It's computation presents challenges.
- Nonetheless it is a powerful tool for exploring high-dimensional probability distributions.
- Using the WD, the ensemble method, and ideas from sampling theory we have designed an experiment to test effects of rounding error on model climatology.
- Half-precision with stochastic-rounding is a suitable arithmetic for climate modelling with both of the L63 and Shallow Water models investigated so far.



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Thank-you!!!:)

... Any questions/thoughts/suggestions?



